Gravitational Acceleration

The gravitational acceleration g in the ocean can be taken to be the following function of latitude ϕ and sea pressure p, or height z relative to the geoid,

$$g/(\text{m s}^{-2}) = 9.780\ 327 \left(1 + 5.3024 x 10^{-3} \sin^2 \phi - 5.8 x 10^{-6} \sin^2 2\phi\right) \left(1 - 2.26 x 10^{-7} z/(\text{m})\right)$$

= 9.780\ 327 \left(1 + 5.2792 x 10^{-3} \sin^2 \phi + 2.32 x 10^{-5} \sin^4 \phi\right) \left(1 - 2.26 x 10^{-7} z/(\text{m})\right) (D.3)
\approx 9.780\ 327 \left(1 + 5.2792 x 10^{-3} \sin^2 \phi + 2.32 x 10^{-5} \sin^4 \phi\right) \left(1 + 2.22 x 10^{-7} z/(\text{m})\right).

The dependence on latitude in Eqn. (D.3) is from Moritz (2000) and is the gravitational acceleration on the surface of an ellipsoid which approximates the geoid. The variation of g with z and p in the ocean in Eqn. (D.3) is derived in McDougall *et al.* (2010b). The height z above the geoid is negative in the ocean. Note that g increases with depth in the ocean at about 71.85% of the rate at which it decreases with height in the atmosphere.

At a latitude of 45°N and at p = 0, $g = 9.8062 \text{ m s}^{-2}$, which is a value commonly used in ocean models. The value of g averaged over the earth's surface is $g = 9.7976 \text{ m s}^{-2}$, while the value averaged over the surface of the ocean is $g = 9.7963 \text{ m s}^{-2}$ (Griffies (2004)).